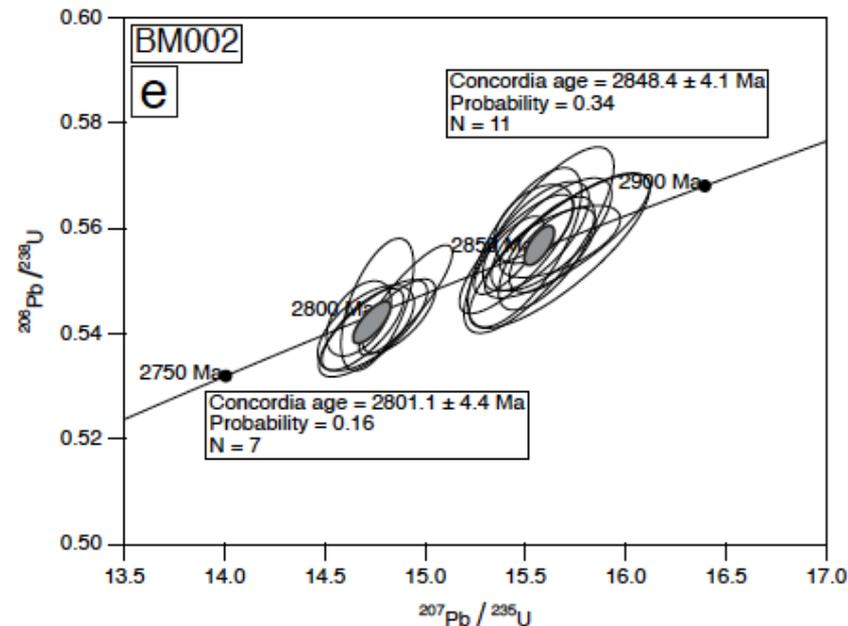
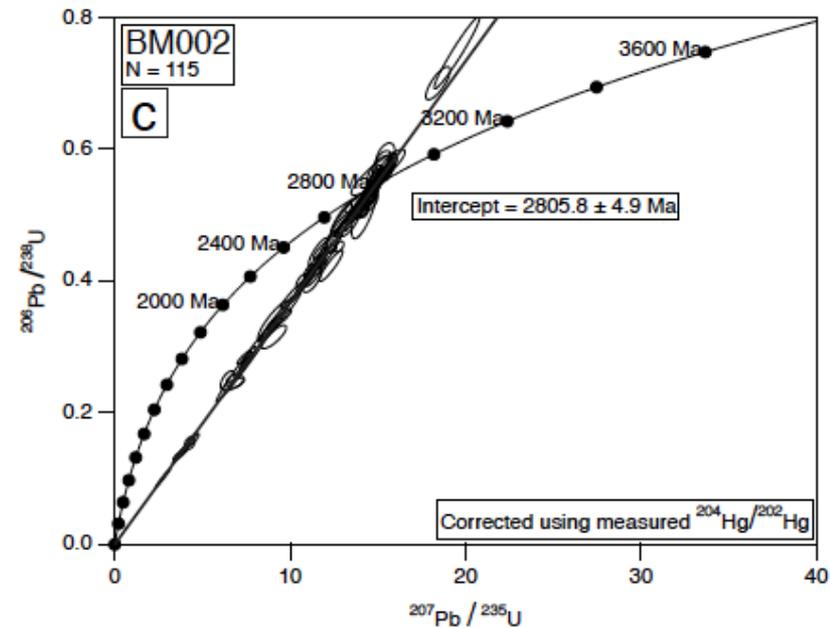


Correcting for common Pb in standards

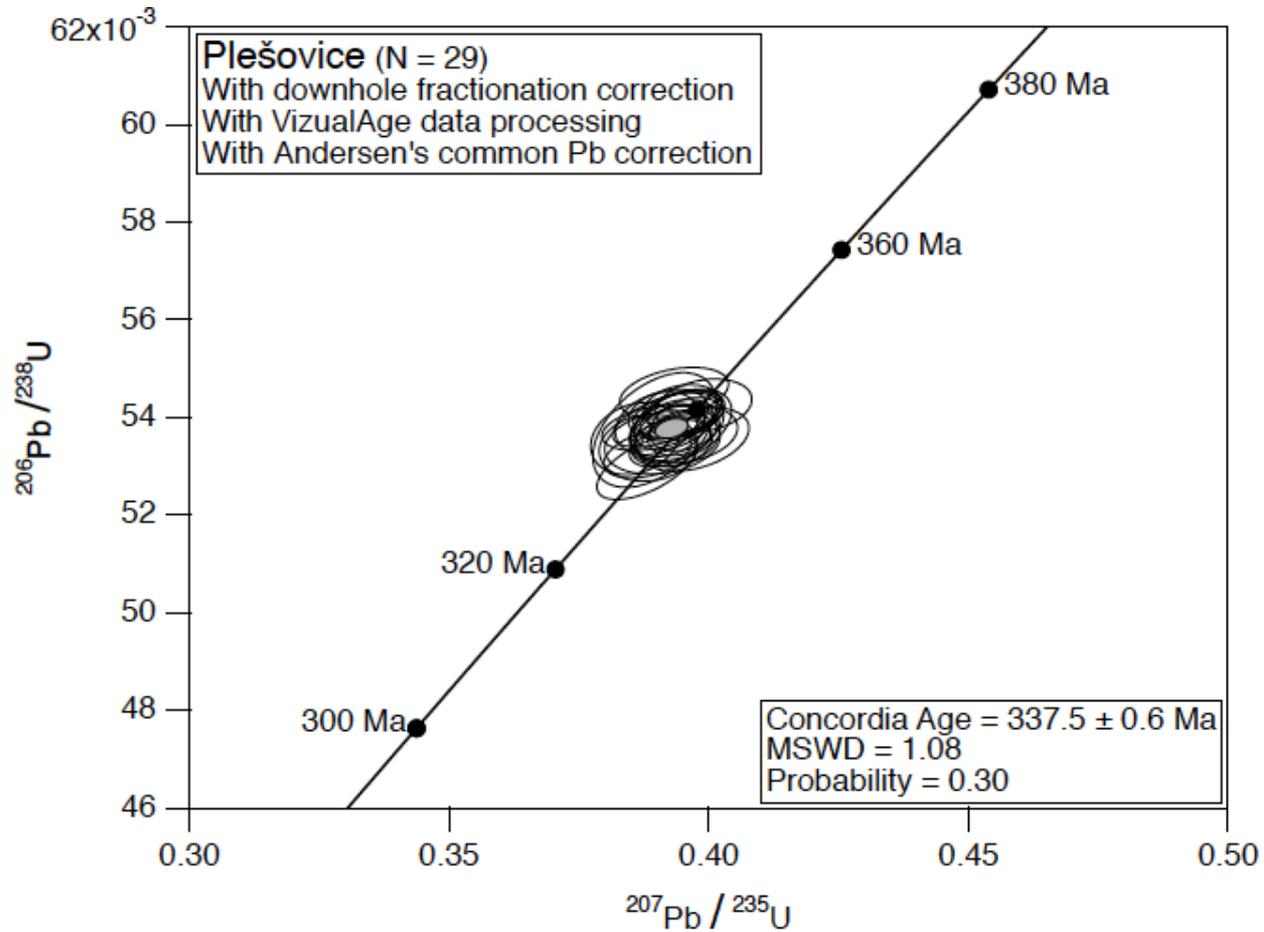
- ▶ A brief summary of VizualAge_UcomPbine (Chew, Petrus & Kamber, Chemical Geology 2014)
- ▶ Corrects for variable common Pb in standards (using either a ^{204}Pb -, ^{207}Pb - or ^{208}Pb correction) prior to correcting for LIEF and session drift
- ▶ It assumes:
 - ▶ 1) standards are age homogenous if they didn't contain common Pb;
 - ▶ 2) the “**end member**” common Pb is isotopically homogenous
 - ▶ 3) However there can be **variable incorporation of the amount of common Pb** – either from standard grain to grain, or even variable amounts of common Pb during an individual TRA standard grain analysis

VizualAge (Petrus & Kamber, 2012)

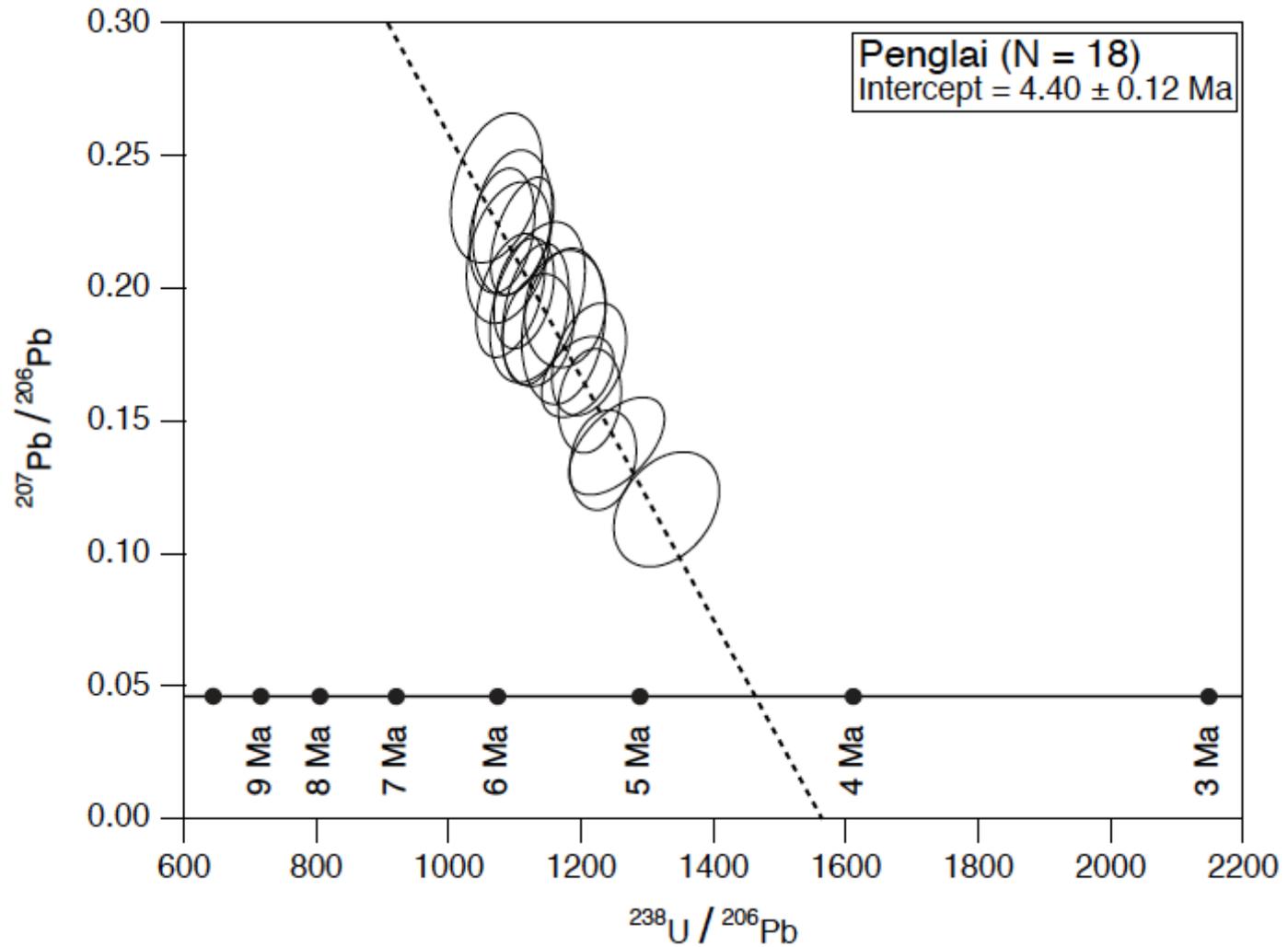
- ▶ Data reduction scheme for Iolite
- ▶ $^{207}\text{Pb}/^{206}\text{Pb}$ dates;
- ▶ 'live' concordia;
- ▶ 'live' error ellipses;
- ▶ ^{204}Pb common Pb to unknowns



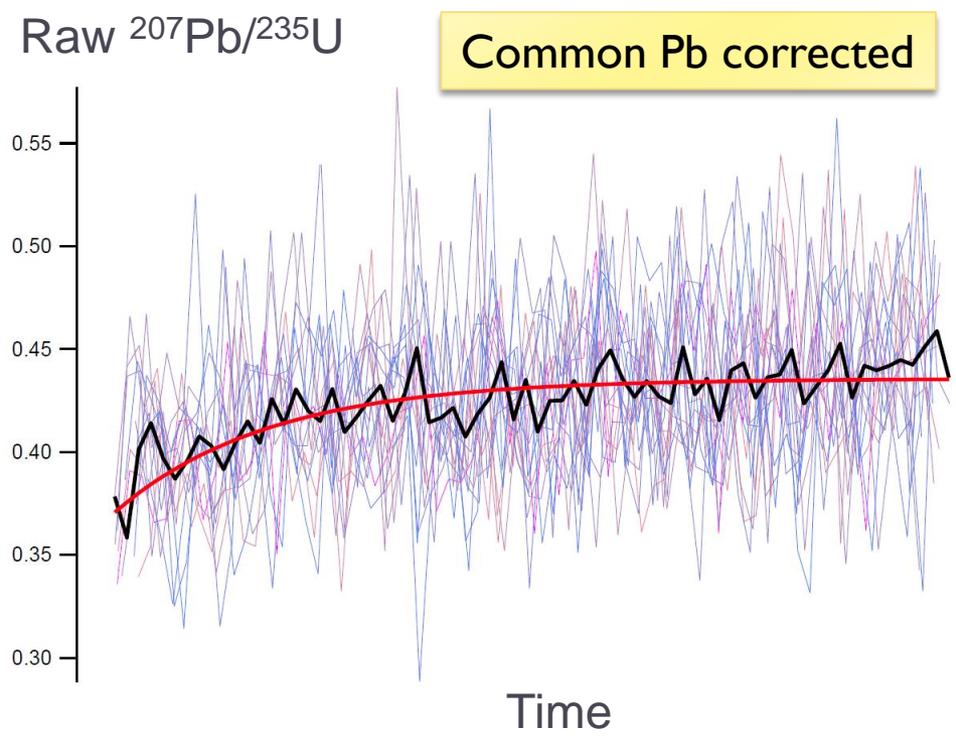
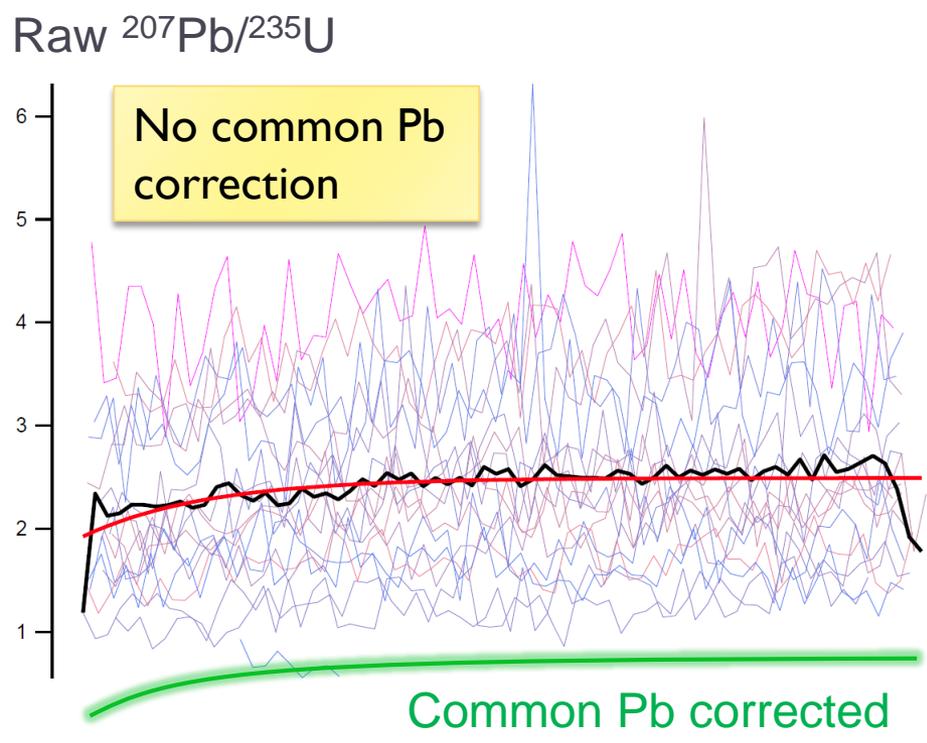
VizualAge example I



VizualAge example 2



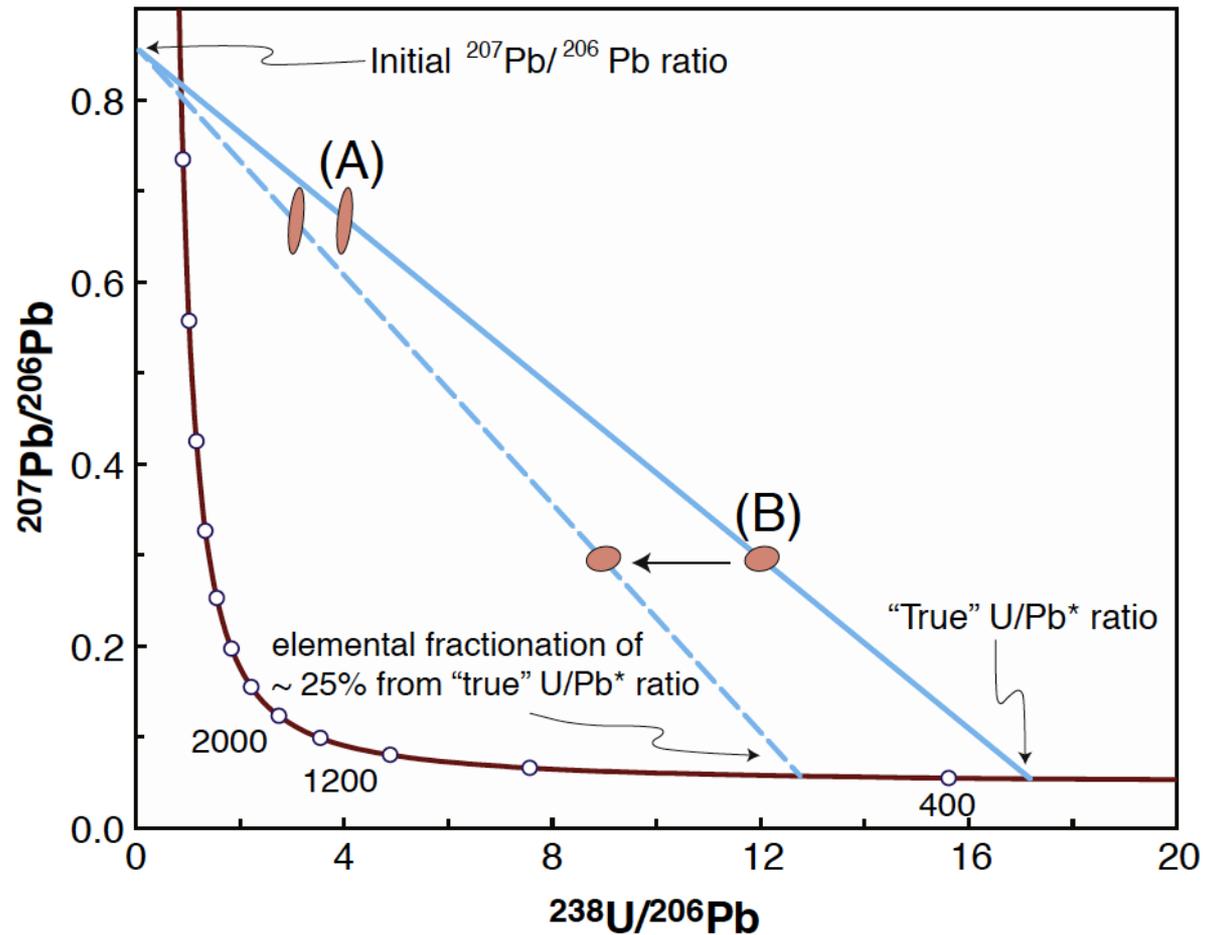
The problem of common Pb in standards



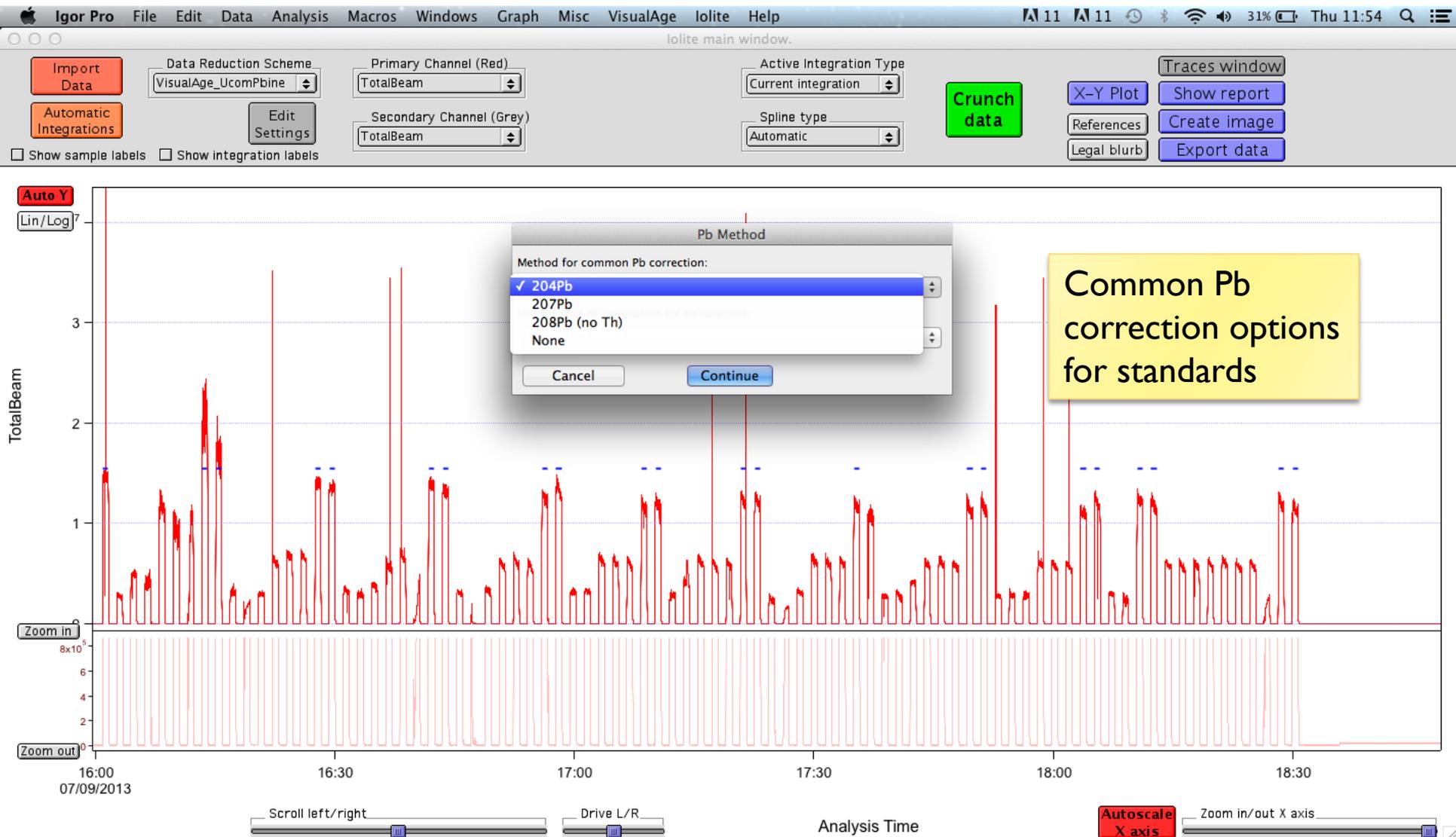
Standard: c. 523.5 Ma McClure Mountain apatite

VizualAge_UcomPbine

- ▶ Assume Pb isotopic ratios are essentially unaffected by LIEF
- ▶ Correct standards for common Pb prior to downhole fractionation correction
- ▶ Deviation from “true” U/Pb ratio is due to elemental fractionation
- ▶ Correct for this by sample-standard bracketing

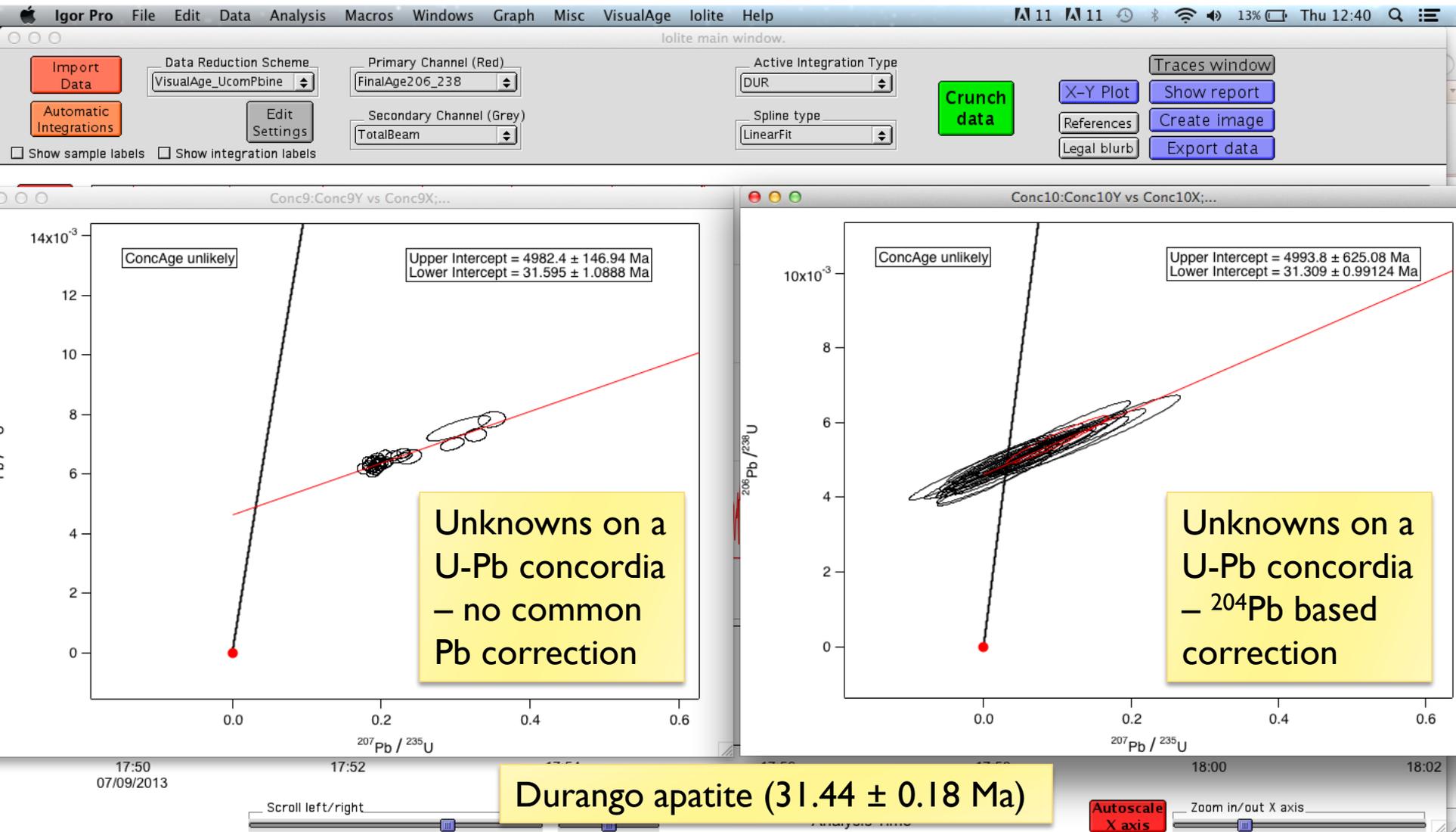


VizualAge_UcomPbine

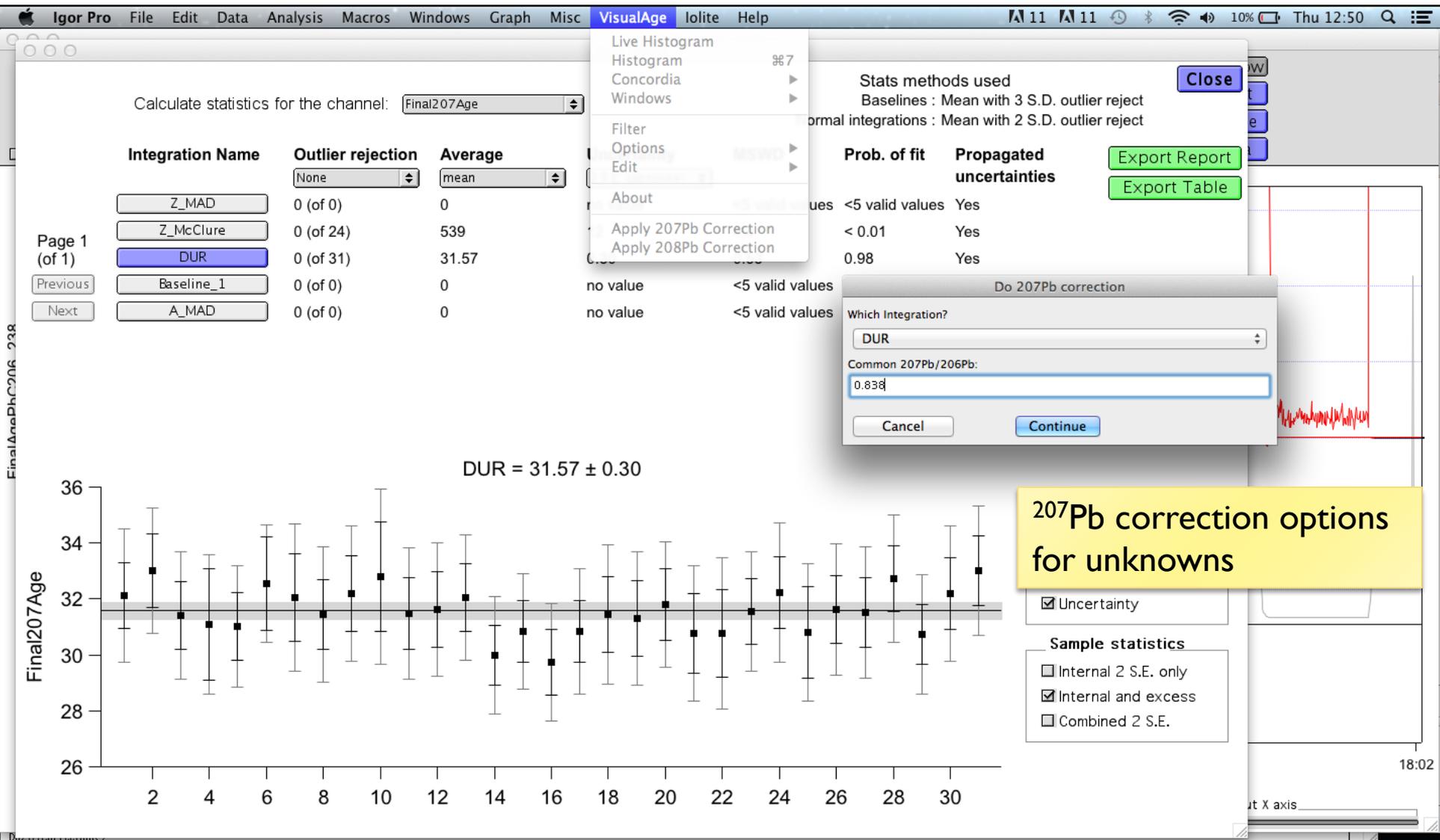


Chew et al. (2014), Chemical Geology

VizualAge_UcomPbine



VizualAge_UcomPbine



^{207}Pb correction options for unknowns

- Uncertainty
- Sample statistics**
- Internal 2 S.E. only
- Internal and excess
- Combined 2 S.E.

VizualAge_UcomPbine: summary

COMMON Pb CORRECTION TO STANDARDS:

- ▶ 3 methods: ^{204}Pb -, ^{207}Pb - and ^{208}Pb -correction

COMMON Pb CORRECTION TO UNKNOWNNS:

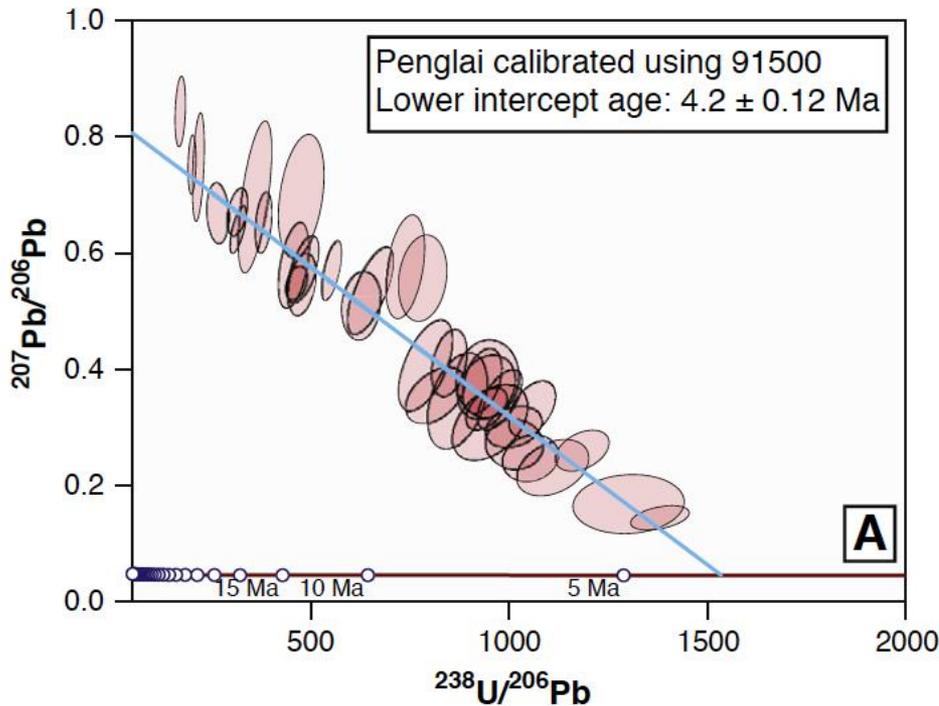
- ▶ ^{204}Pb -, ^{207}Pb - and ^{208}Pb -correction. ^{204}Pb method uses conventional VizualAge correction; ^{207}Pb - and ^{208}Pb -correction user inputs initial Pb ratio

CONCORDIA OPTIONS:

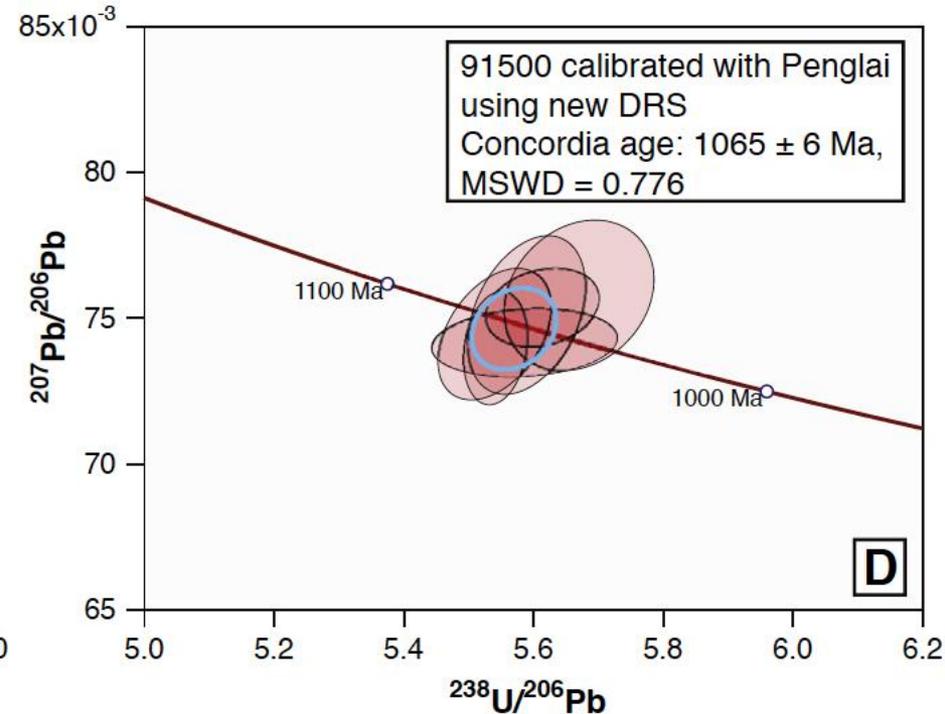
- ▶ Live Concordias; Wetherill and Tera-Wasserburg: non-corrected or ^{204}Pb and ^{208}Pb -correction

VizualAge_UcomPbine: nice example!

30 micron zircon
spot analyses



This is a common Pb-infested Penglai zircon (4.1 Ma), with some analyses plotting close to modern day common Pb. 91500 used as the primary.



Same session – but we used the common Pb infested Penglai as the primary and treated 91500 as the unknown – comes out at 1065 Ma.

1. Application-specific strategies

Listed in (a crude) order of increasing common Pb

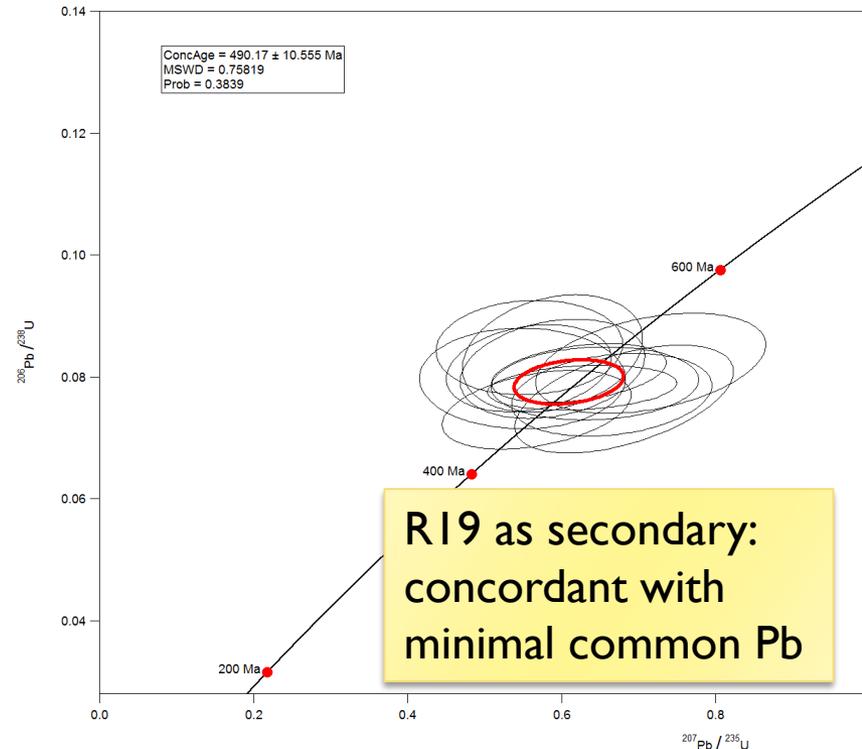
- ▶ 1. Rutile
 - ▶ 2. Titanite
 - ▶ 3. Apatite
 - ▶ 4. Calcite
-
- ▶ In Trinity College Dublin, we use a Photon Machines Analyte Exite ArF Excimer laser coupled to a Thermo Scientific iCAP Qc
 - ▶ For rutile, apatite and titanite, a spot size of 30 to 60 microns (depending on the U and Pb contents in the session), 5Hz rep rate, 45 sec ablation, 25 sec background, 1 primary and 2 secondaries (blocks of 6 standards then 20 unknowns)
 - ▶ All standards corrected for common Pb prior to LIEF and sample-standard bracketing using a modified version of Vizual Age (VizualAge_UcomPbine)

Rutile

▶ Main standards:

- ▶ **R10/R10b (Luvizotto et al. 2009)**
- ▶ R19 (Zack et al., 2011)
- ▶ Both standards contain minimal common Pb and are typically concordant

Primary



- ▶ Rutile ideally suited to a ^{208}Pb correction due to low Th
- ▶ If no Th present, all ^{208}Pb assumed common
- ▶ Some Th can be present in unknowns. As standards contain no Th how do we determine $^{232}\text{Th}/^{208}\text{Pb}$ fractionation? Tune on NIST with Th/U~1 and measure $^{232}\text{Th}/^{238}\text{U}$ of NIST during the session?
- ▶ So can do ^{204}Pb , ^{207}Pb and ^{208}Pb corrections and compare

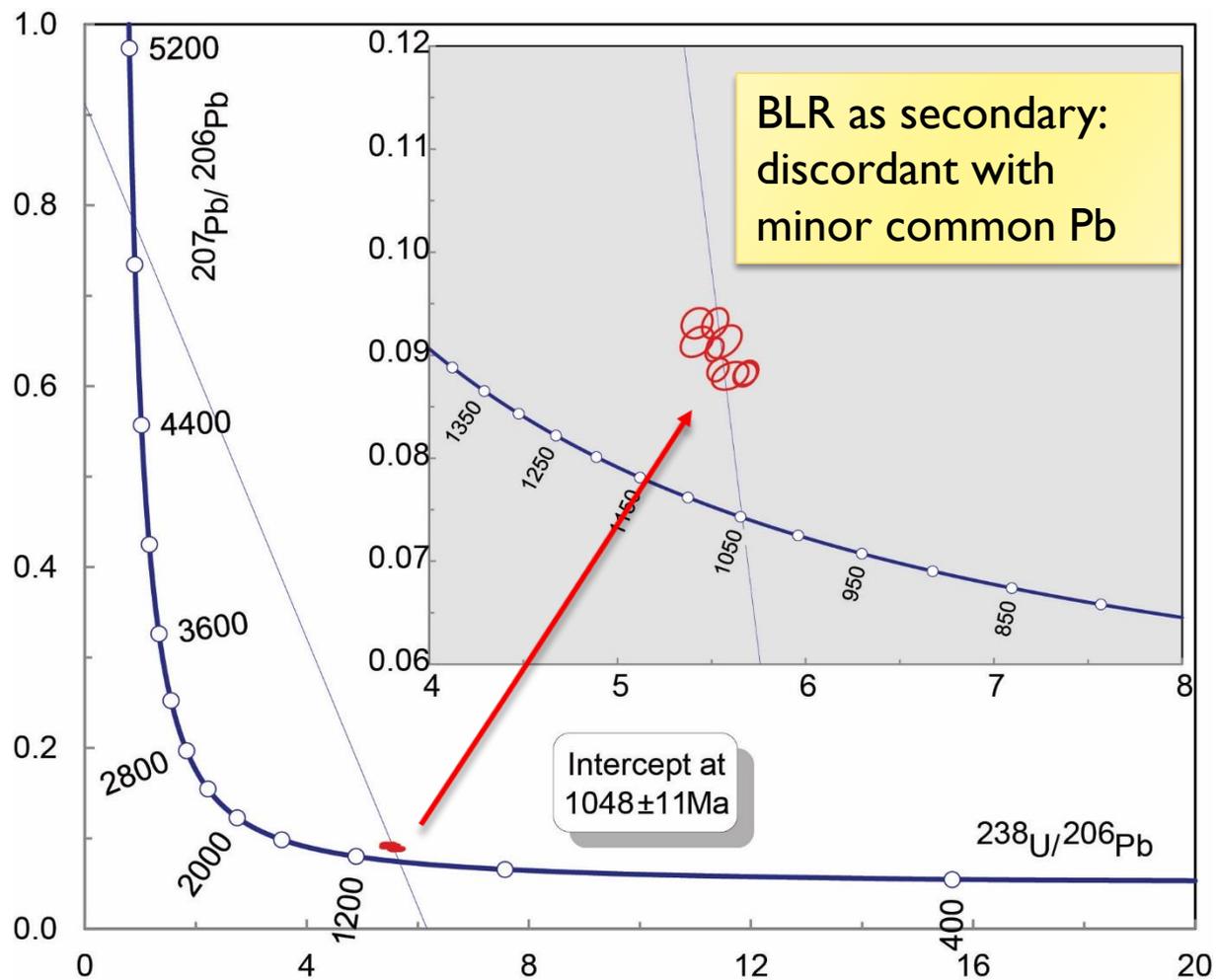
Titanite

Primary



- ▶ Some crystal standards:
 - ▶ OLT₁ (Kennedy et al., 2010)
 - ▶ BLR (Aleinikoff et al., 2007; UCSB group)
 - ▶ Khan (Heaman et al., 2009)
 - ▶ These are large crystal standards that contain minor common Pb- typically minor discordance but analyses overlap
- ▶ Mineral separates
 - ▶ Fish Canyon tuff
 - ▶ McClure Mountain syenite (Schone and Bowring, 2006)
 - ▶ Variable common Pb from grain to grain

Titanite



- ▶ We do correct standards for common Pb (can get variations of 1 to 2% in U/Pb ratio due to variable common Pb). Our chips of Khan seem to suffer from minor Pb loss.

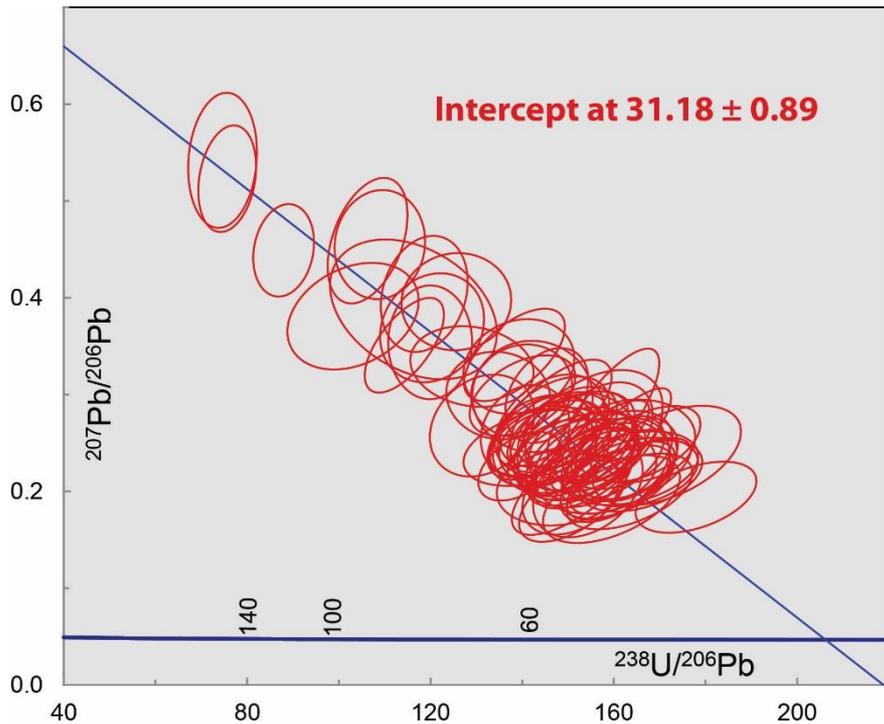
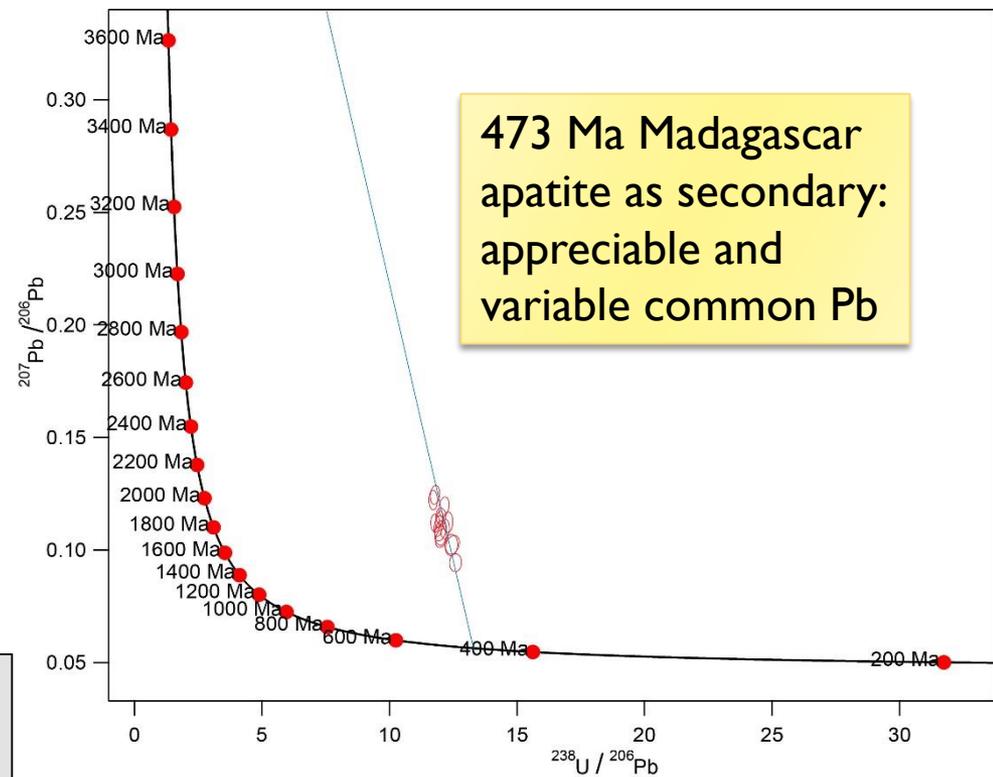
Apatite

Primary



- ▶ Crystal standards:
 - ▶ Madagascar (Thomson et al., 2012)
 - ▶ Durango (McDowell et al., 2005)
 - ▶ These are large crystal standards that contain minor to appreciable common Pb
 - ▶ Standard analyses require common Pb correction
- ▶ Mineral separates
 - ▶ McClure Mountain syenite (Schoene and Bowring, 2006)
 - ▶ Variable common Pb from grain to grain

Apatite

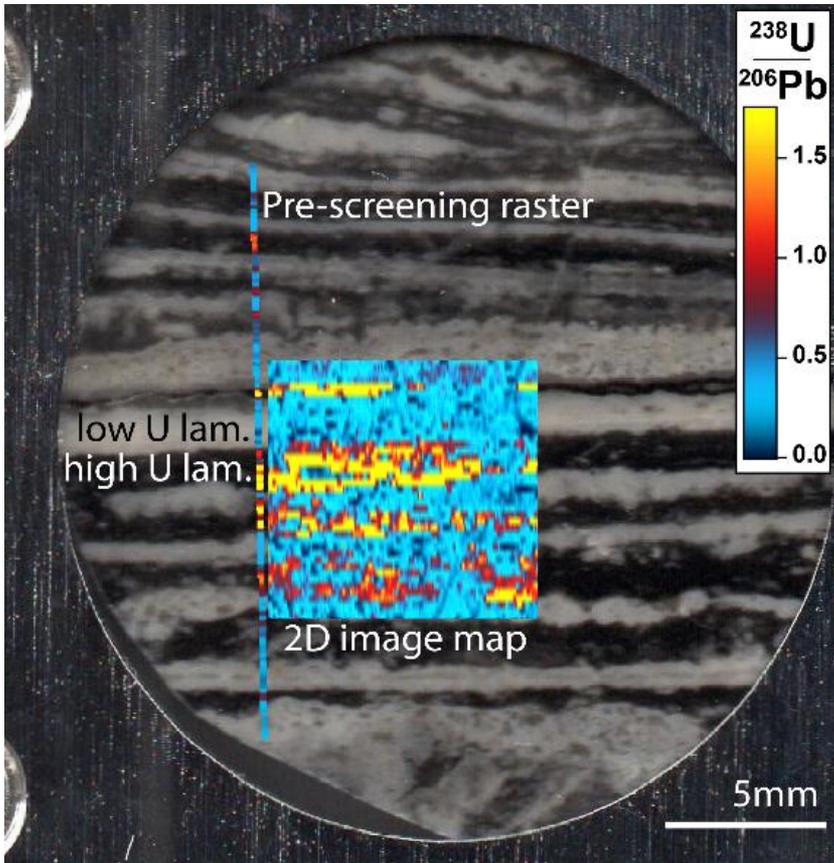


31.44 Ma Durango apatite as secondary: appreciable and variable common Pb

Calcite

- ▶ Age standards:
 - ▶ “Troy” speleothem calcite (Li et al., 2014 Chemical Geology) and that is it....
- ▶ Very tricky – often very low U (100 ppb – 1ppm), lots of common Pb.
- ▶ Li et al. (2014) used LA-MC-ICPMS spot analyses with NIST 614 glass and Troy Calcite (TIMS age of 251 ± 2 Ma) as SRMs
- ▶ Alternative approach described here is image age mapping (“rastering”) by LA-Q-IPCMS.
- ▶ It can often circumvent the problems of low U contents and/or high initial Pb by identifying zones of high U high U/Pb ratio on LA-ICPMS image maps. U-Pb ages are generated from these same image maps. Same standards used as Li et al. (2014)

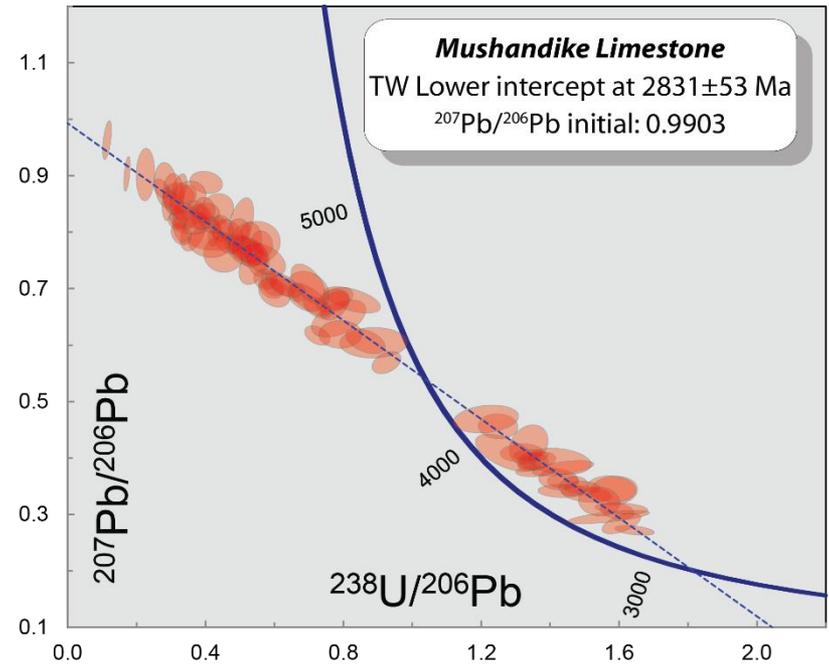
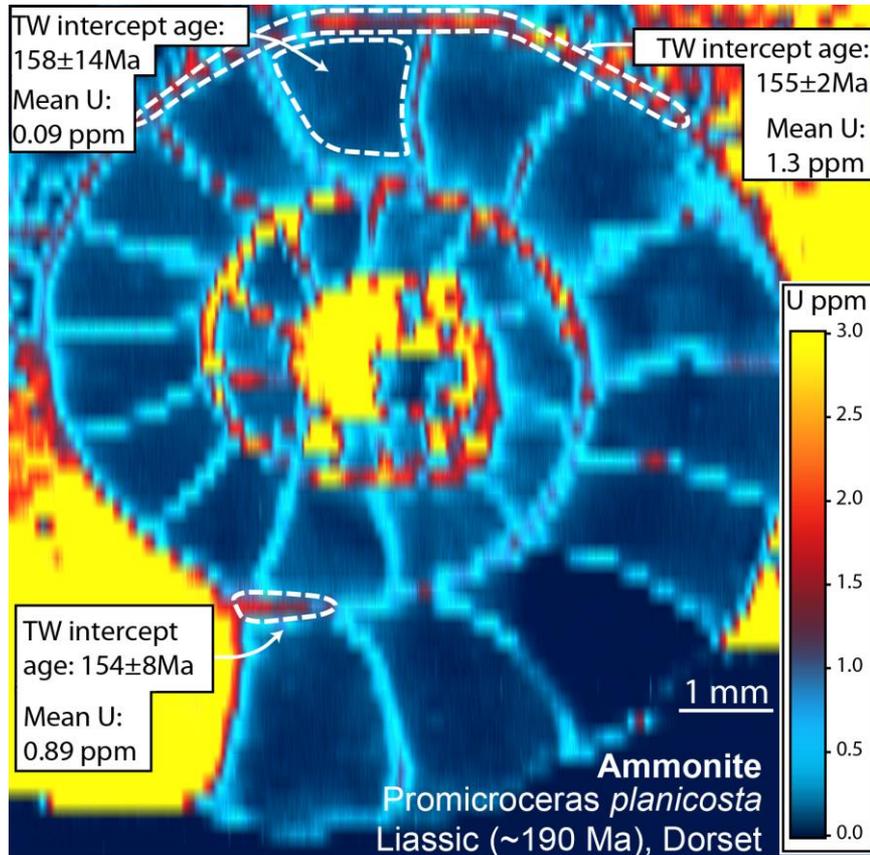
Calcite



LA-ICPMS $^{238}\text{U}/^{206}\text{Pb}$ map of a laminated Neoproterozoic microbial dolomite illustrating dark laminae with high $^{238}\text{U}/^{206}\text{Pb}$ ratios. The image map is overlain on a scanned sample image using IOLITE

- ▶ Rationale - no way to know in advance of analysis if a sample has a viable U/Pb ratio, so many carbonates are not datable.
- ▶ Do a pre-screening raster to identify highly radiogenic subzones (low common Pb and/or high U) that are the key to precise ages.
- ▶ Then map highly radiogenic subzones and generate U-Pb ages from these maps

Calcite – example data



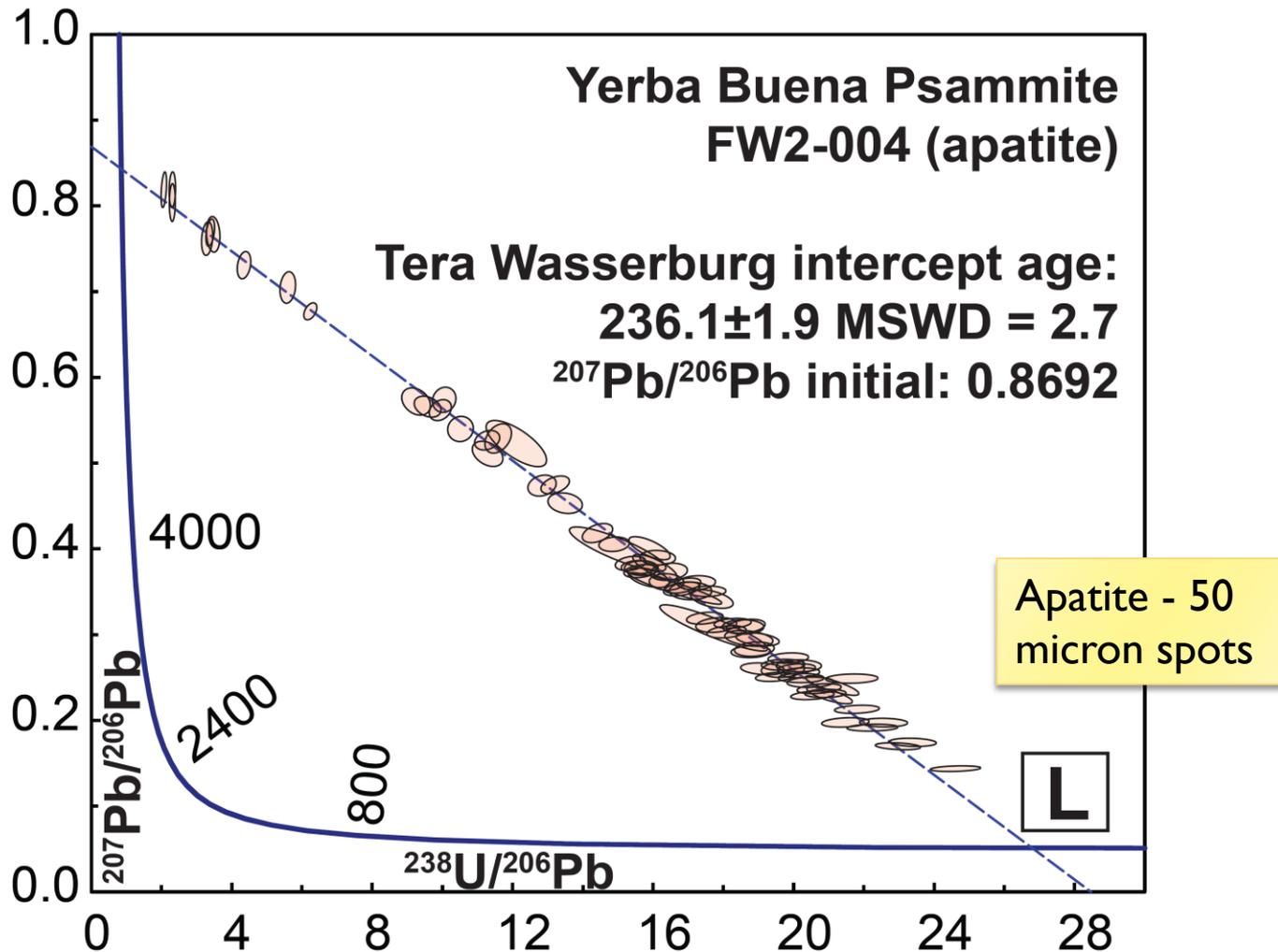
LA-ICPMS U-Pb calcite dating of the Mushandike limestone (2839 ± 33 Ma TIMS age; Moorbath et al., 1987).

LA-ICPMS U concentration map of a diagenetic calcite cement in a Liassic ammonite with U-Pb calcite ages marked

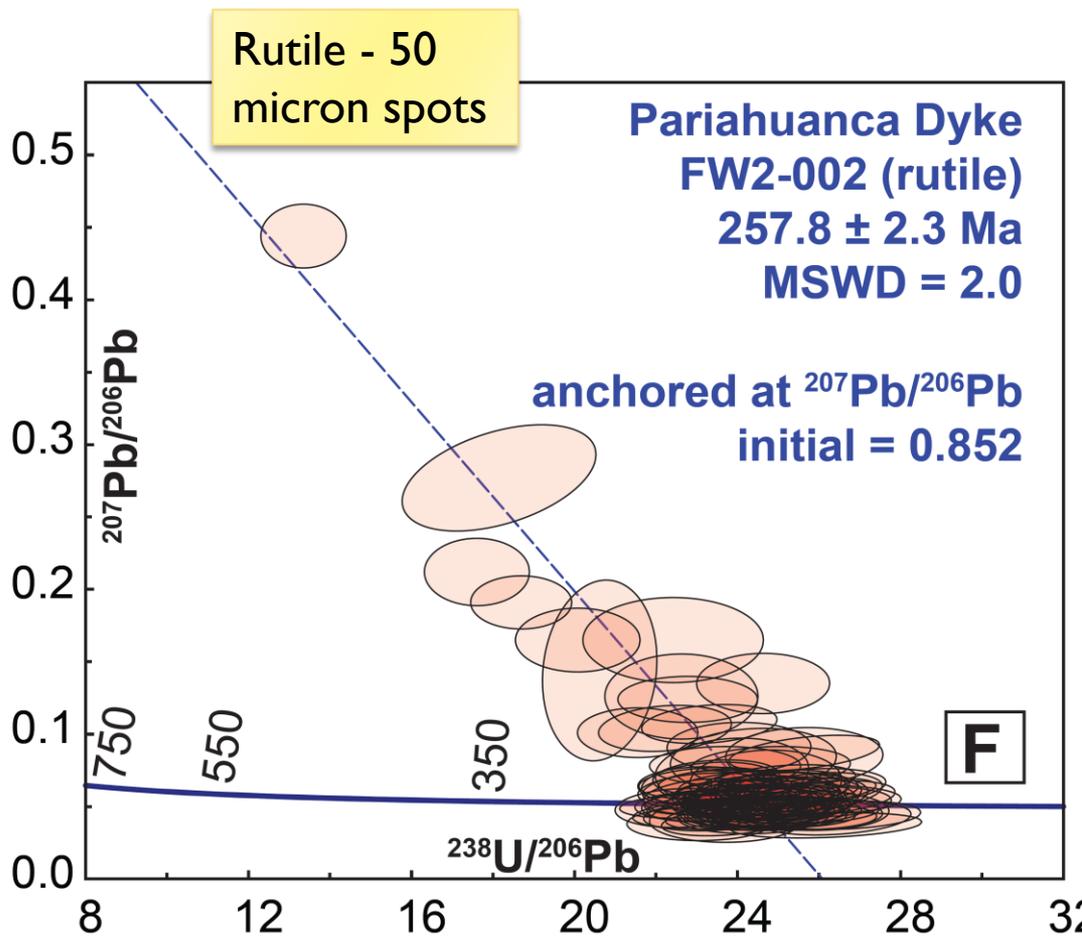
2. Tera-Wasserburg discordia examples

- ▶ CASE A. Well constrained – no anchor required. Data alone define intercept age.
- ▶ CASE B. Moderately constrained, close to Concordia. Data alone define intercept age but a very conservative choice of common Pb anchor improves precision
- ▶ CASE C. Poorly constrained and plot close to common Pb intercept. Data alone should define intercept age and anchoring with initial Pb should not be employed.
- ▶ CASE D. Poorly constrained intercept (analyses cluster with no spread) but close to Concordia. A very conservative choice of common Pb anchor should be used.

CASE A: Data alone define intercept



CASE A: Data alone define intercept



**TW Intercept age anchored at
 $^{207}\text{Pb}/^{206}\text{Pbc} = 0.83 \pm 0.02$**

257.3 ± 2.3 Ma, MSWD = 2

**TW Intercept age anchored at
 $^{207}\text{Pb}/^{206}\text{Pbc} = 0.874 \pm 0.02$**

257.9 ± 2.3 Ma, MSWD = 2

**Unanchored TW Intercept age
 257.8 ± 2.4 Ma, MSWD = 2**

⇒ Large enough spread in data to give a well constrained unanchored intercept

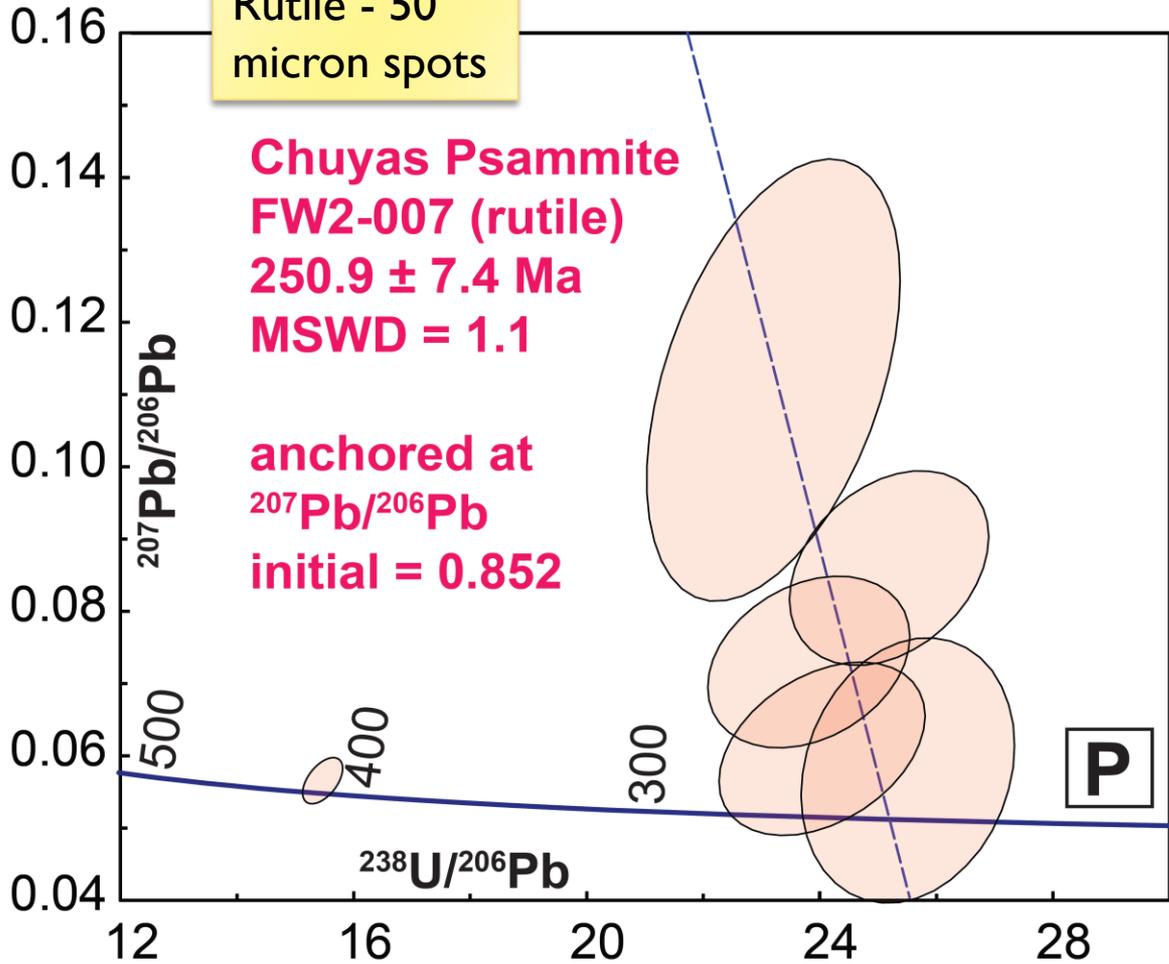
⇒ Unanchored and anchored ages virtually identical

CASE B: Moderately constrained isochron

Rutile - 50
micron spots

**Chuyas Psammite
FW2-007 (rutile)
250.9 ± 7.4 Ma
MSWD = 1.1**

**anchored at
 $^{207}\text{Pb}/^{206}\text{Pb}$
initial = 0.852**



**TW Intercept age anchored at
 $^{207}\text{Pb}/^{206}\text{Pbc} = 0.83 \pm 0.02$**

250.7 ± 7.4 Ma, MSWD = 1.1

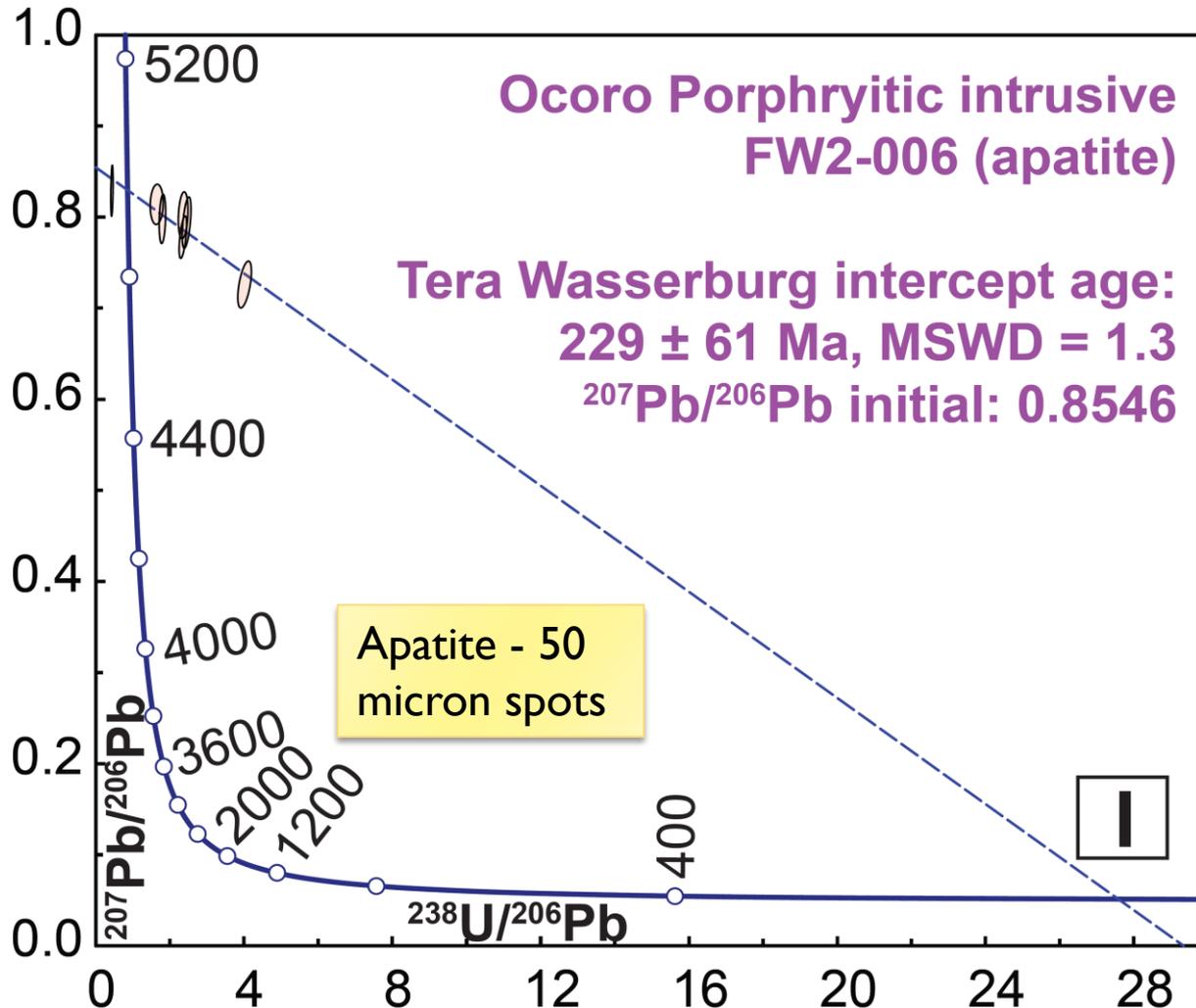
**TW Intercept age anchored at
 $^{207}\text{Pb}/^{206}\text{Pbc} = 0.874 \pm 0.02$**

251.1 ± 7.4 Ma, MSWD = 1.1

**Unanchored TW Intercept age
 252.0 ± 12 Ma, MSWD = 1.4**

⇒ Would seem reasonable to anchor isochron at Stacey & Kramers with large uncertainty – not enough spread in data to give a well constrained unanchored intercept

CASE C: Poorly defined intercept; data plot close to common Pb



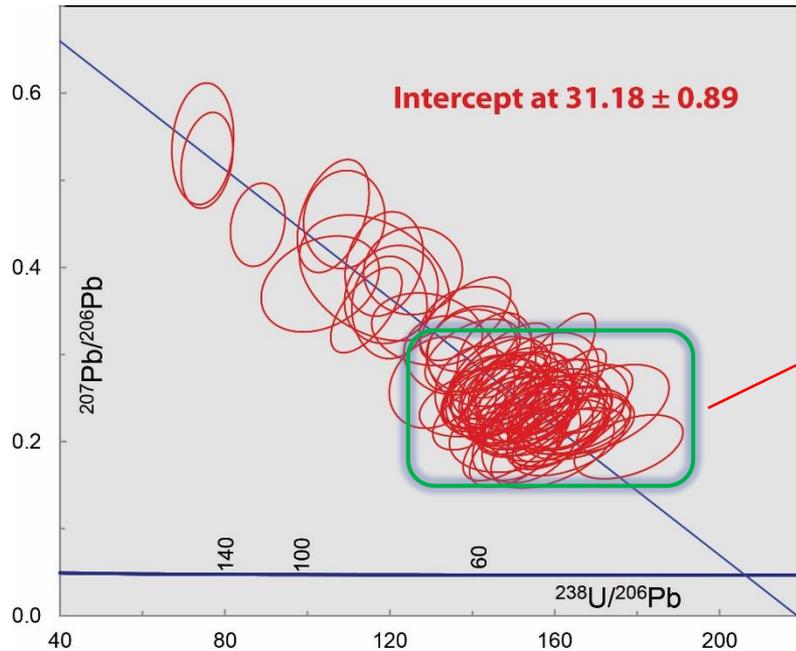
**TW Intercept age using
 $^{207}\text{Pb}/^{206}\text{Pb}_c = 0.85 \pm 0.02$
227 ± 57 Ma, MSWD = 1.1**

**Unanchored Intercept age
229 ± 61 Ma, MSWD = 1.3**

- ⇒ No spread in data to give a well constrained unanchored intercept
- ⇒ Need to employ conservative uncertainty in initial composition of Pb; does not help precision

CASE D: Poorly defined intercept; data plot close to concordia

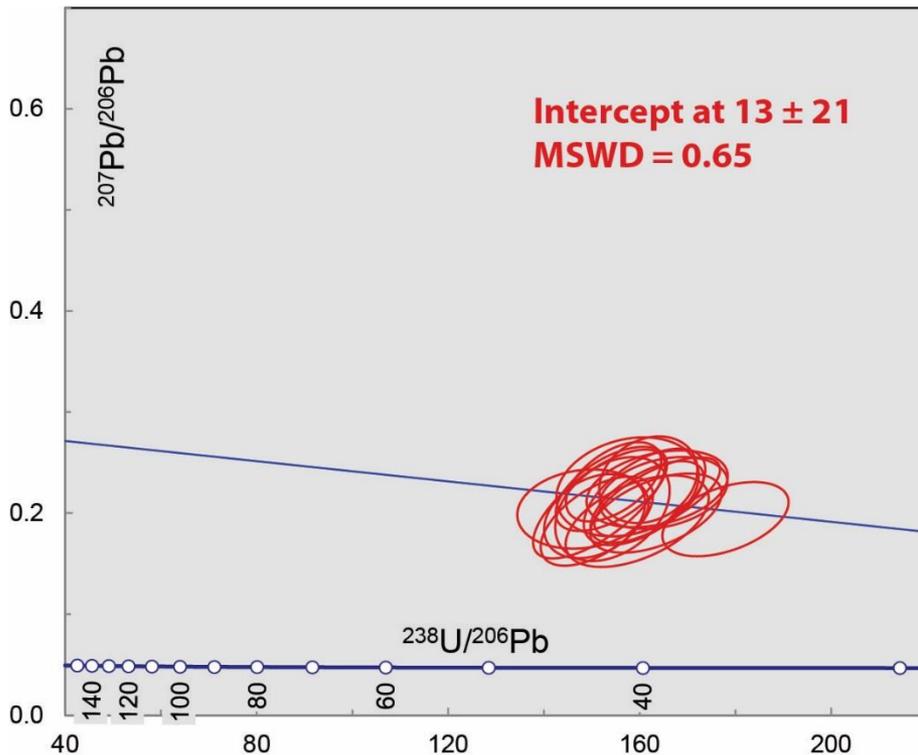
Durango apatite - 50 micron spots
=CASE A (well constrained)



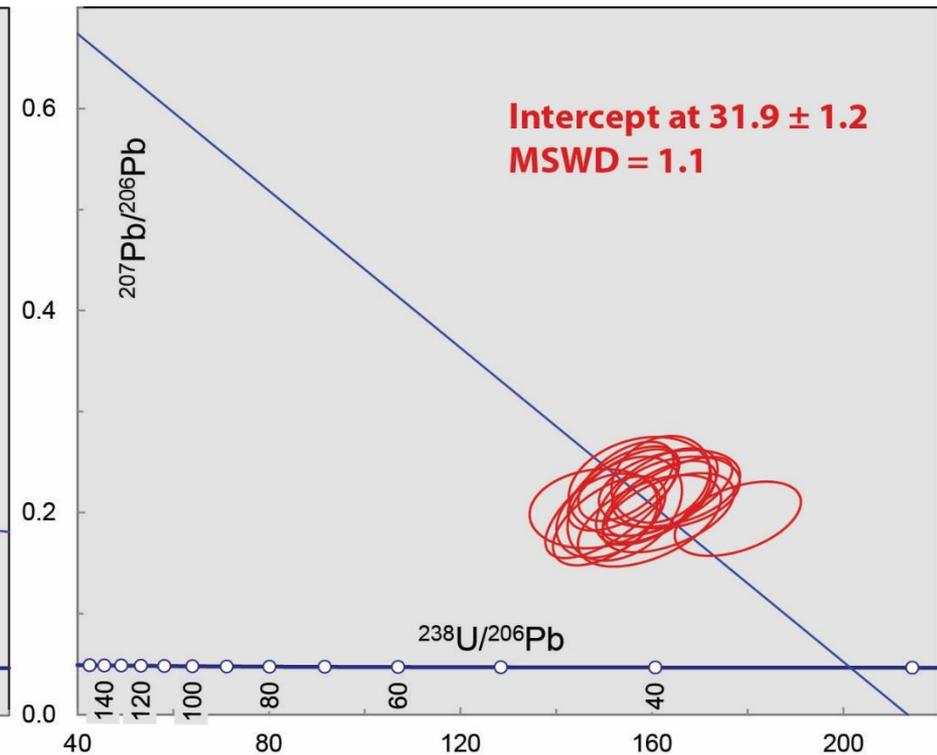
Will take a subset of the data to make an overlapping cluster of points close to concordia

CASE D: Poorly defined intercept; data plot close to concordia

31.44 Ma Durango apatite - 50 micron spots



Data alone define intercept

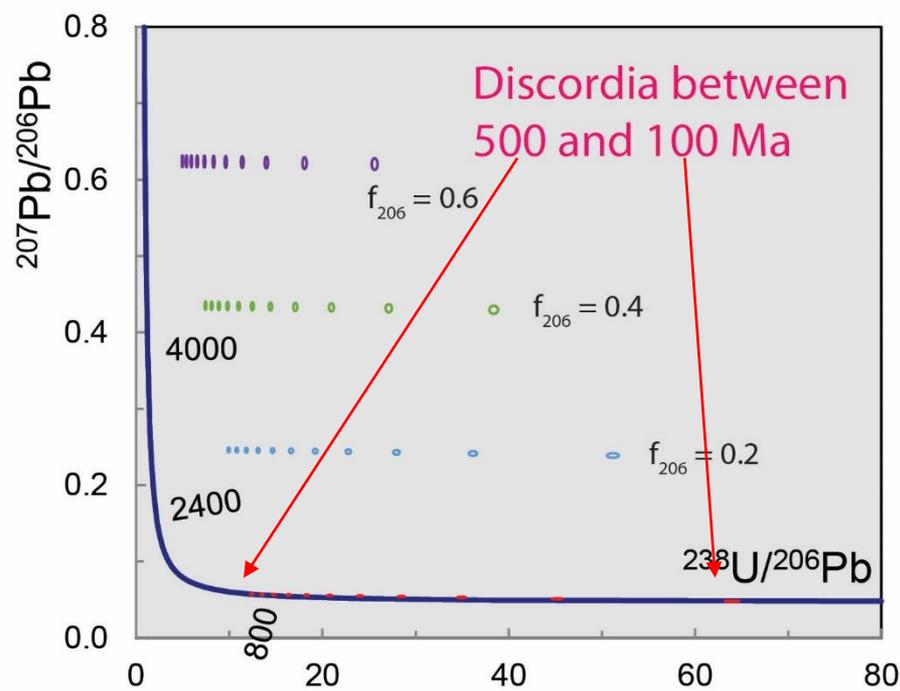
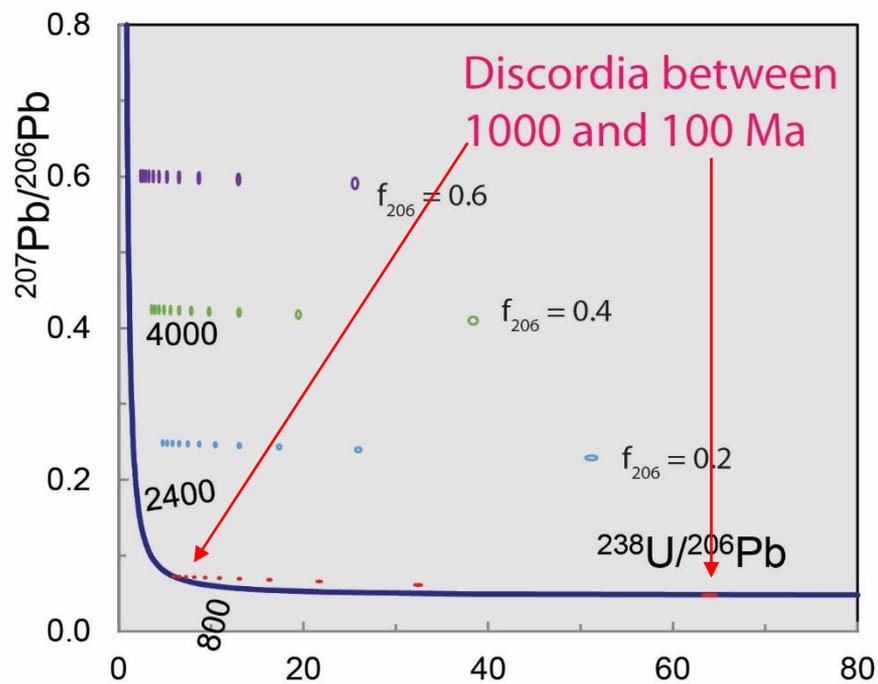
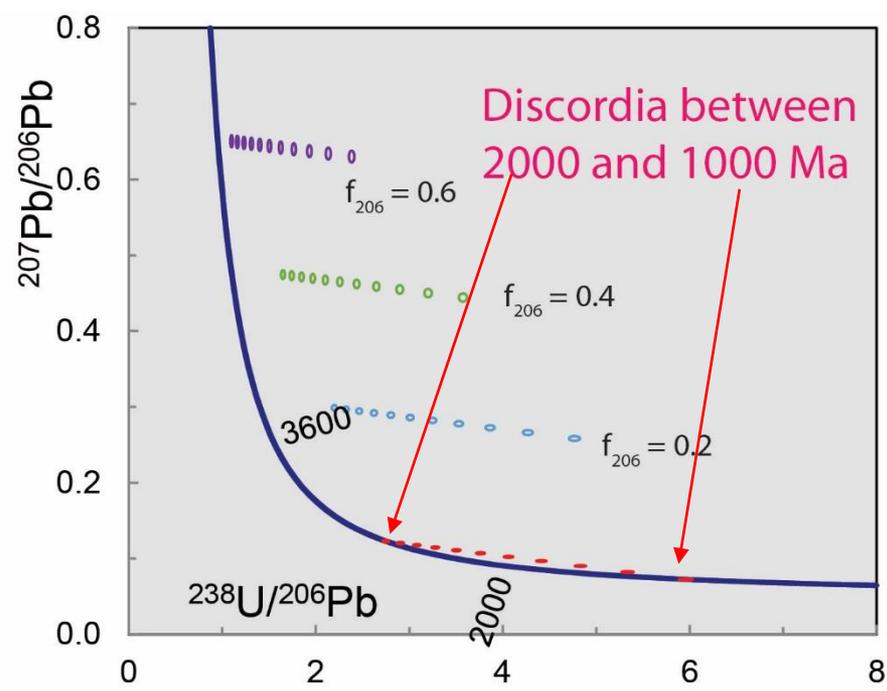
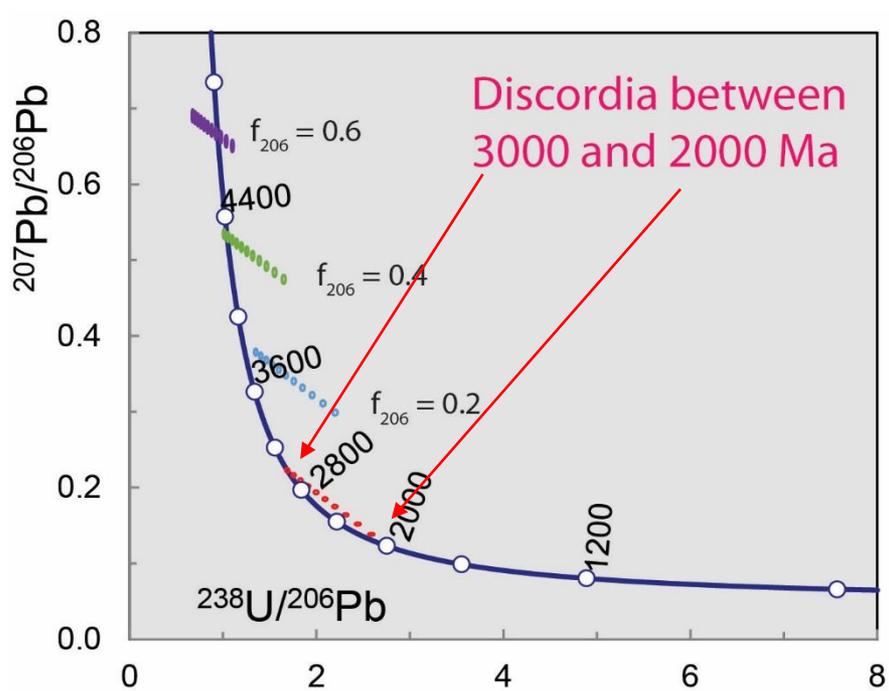


Data anchored at $^{207}\text{Pb}/^{206}\text{Pb}_c = 0.84 \pm 0.03$

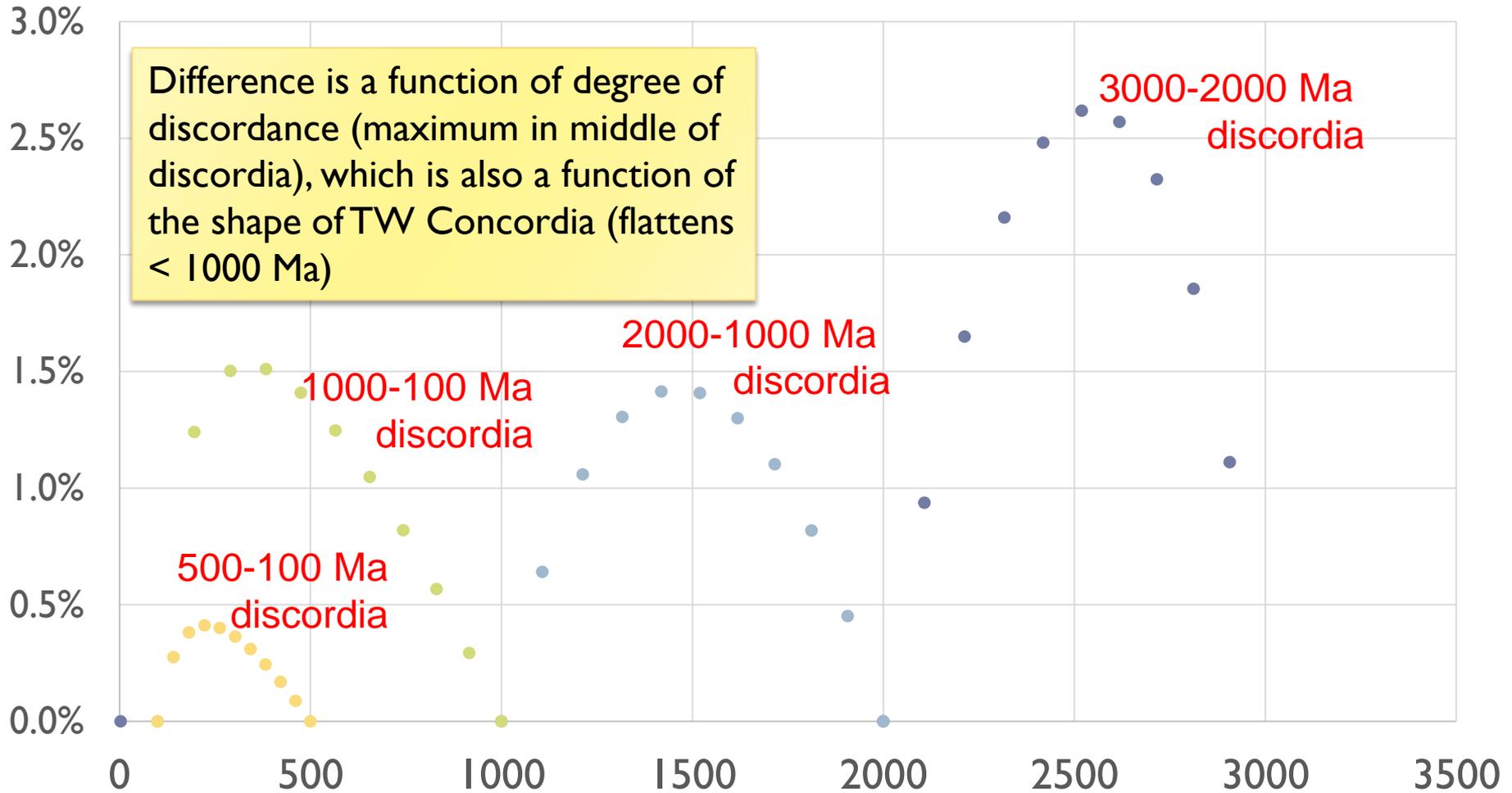
3. When to use ^{204}Pb correction vs ^{207}Pb correction?

[A theoretical discussion independent of analytical setup]

- ▶ A ^{207}Pb correction assumes concordance so should never be used for strongly discordant data (e.g. zircon with significant Pb loss)
- ▶ Rutile, titanite and apatite generally are happier incorporating Pb in their structure (they accept common Pb) and so Pb loss is often less significant
- ▶ But apatite in particular can lose Pb due to slow cooling through the closure temperature window ($T_c = 550 - 375^\circ\text{C}$ depending on grain size and cooling rate)
- ▶ Following diagrams illustrate 4 discordias on TW Concordia (with incorporation of common Pb at $f_{206} = 0.2, 0.4$ and 0.6)
- ▶ Difference in the ^{207}Pb - vs ^{204}Pb -corrected age is then calculated



^{207}Pb corrected vs ^{204}Pb corrected ages (difference %)

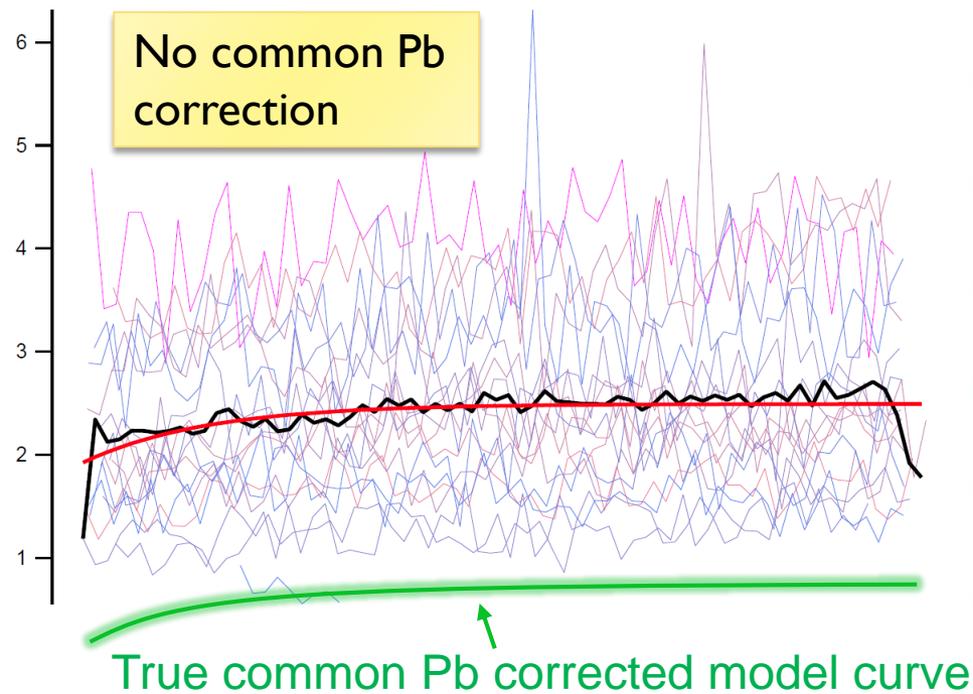


4. Correcting for common Pb in standards

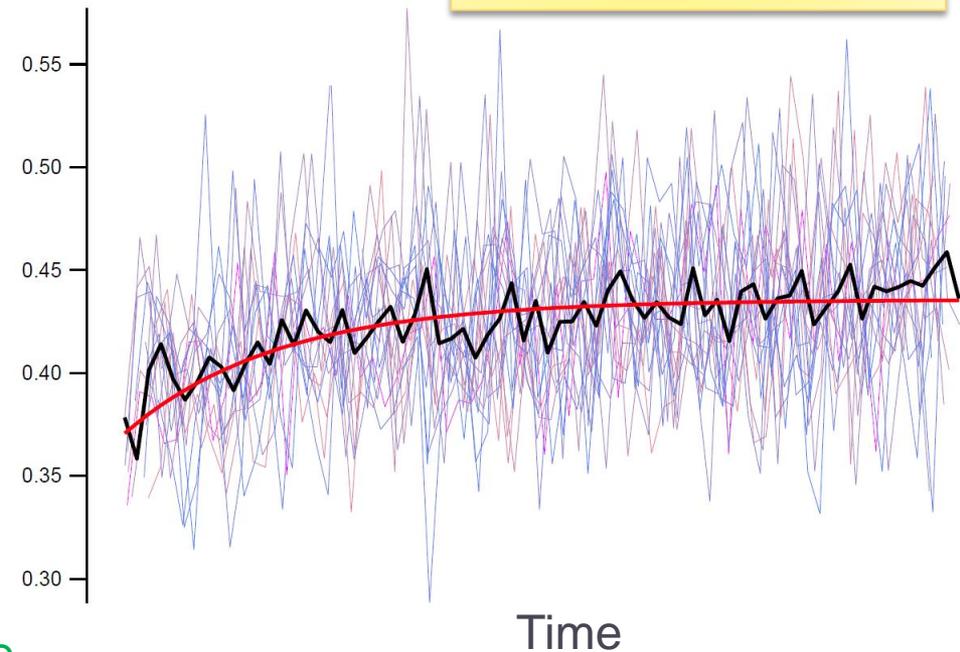
- ▶ A brief summary of VizualAge_UcomPbine (Chew et al. Chemical Geology 2014)
- ▶ Corrects for variable common Pb in standards (using either a ^{204}Pb -, ^{207}Pb - or ^{208}Pb correction) prior to correcting for LIEF and session drift
- ▶ It assumes:
 - ▶ 1) standards are concordant if they didn't contain common Pb;
 - ▶ 2) the "end member" common Pb is isotopically homogenous
 - ▶ 3) However there can be **variable incorporation of the amount of common Pb** – either from standard grain to grain, or even variable amounts of common Pb during an individual TRA standard grain analysis

VizualAge_UcomPbine – correct for common Pb prior to downhole fractionation correction

Raw $^{207}\text{Pb}/^{235}\text{U}$



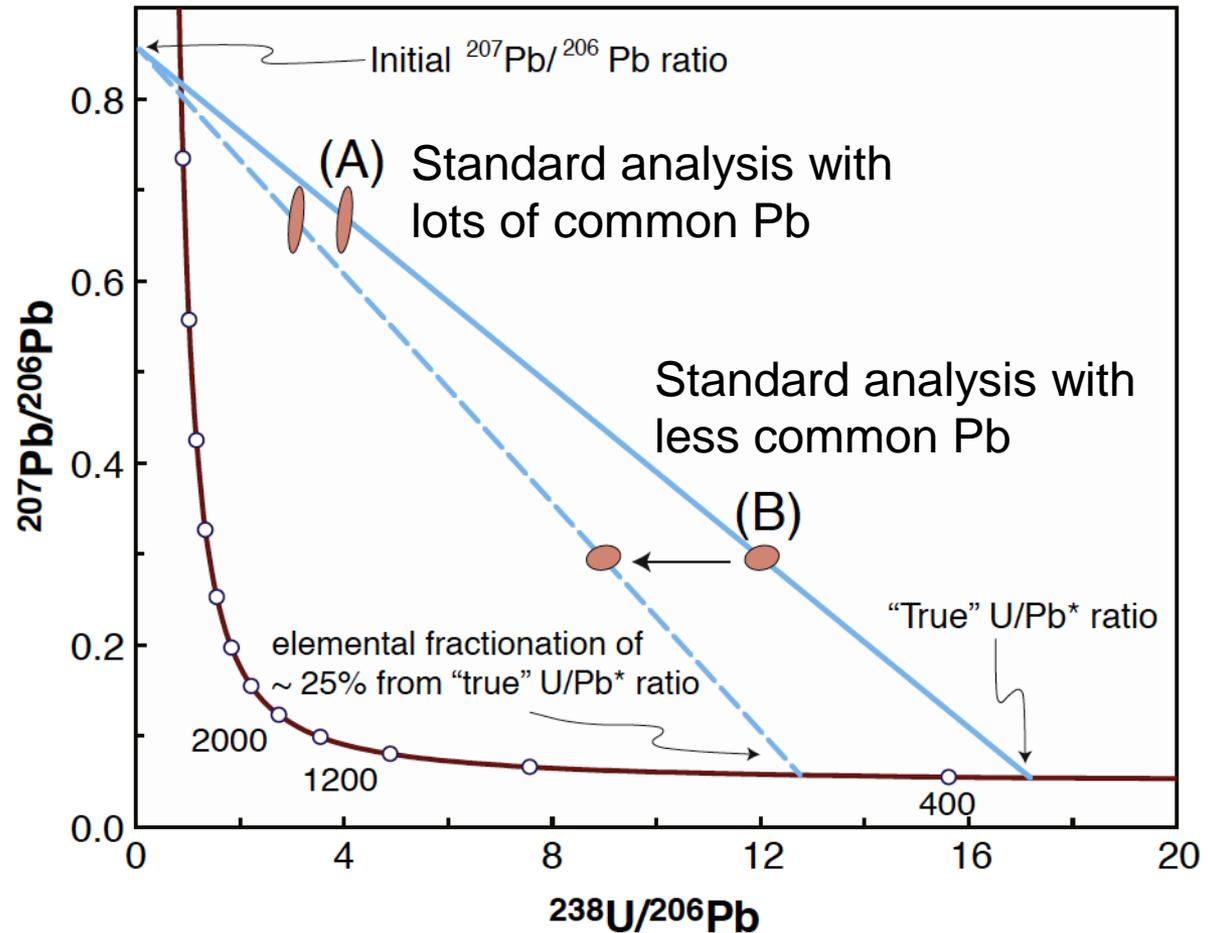
Raw $^{207}\text{Pb}/^{235}\text{U}$



Standard: c. 523.5 Ma McClure Mountain apatite

Correct for session drift by sample-standard bracketing

- ▶ Assume Pb isotopic ratios are unaffected by elemental fractionation
- ▶ Correct standards for common Pb
- ▶ Deviation from "true" U/Pb ratio is due to elemental fractionation
- ▶ Correct for this by sample-standard bracketing



VizualAge_UcomPbine: summary

Common Pb correction to standards :

- ▶ 3 methods: ^{204}Pb -, ^{207}Pb - and ^{208}Pb -correction

Common Pb correction to unknowns:

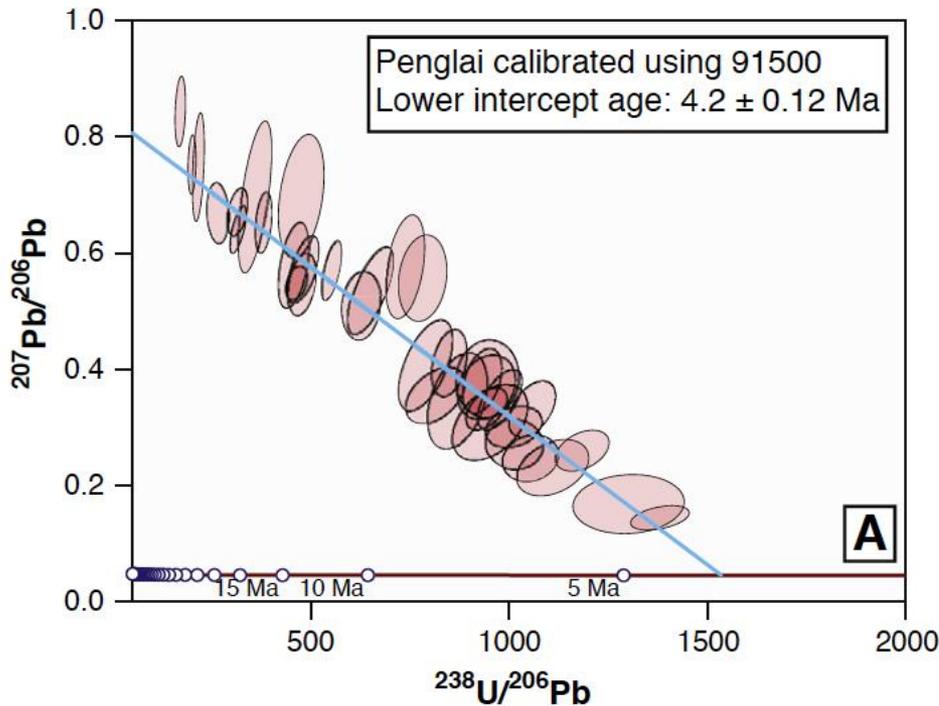
- ▶ ^{204}Pb -, ^{207}Pb - and ^{208}Pb -correction. ^{204}Pb method uses conventional VizualAge correction; ^{207}Pb - and ^{208}Pb -correction user inputs initial Pb ratio

Concordia Options:

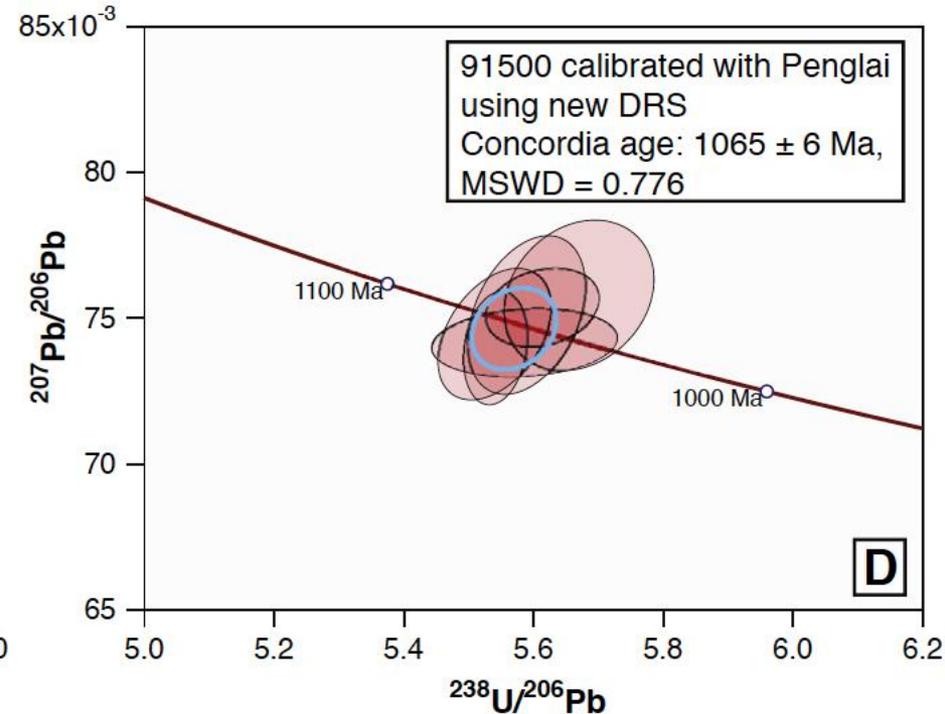
- ▶ Wetherill and Tera-Wasserburg: non-corrected or ^{204}Pb and ^{208}Pb -correction

VizualAge_UcomPbine: nice example!

30 micron zircon
spot analyses



This is a common Pb-infested Penglai zircon (4.1 Ma), with some analyses plotting close to modern day common Pb. 91500 used as the primary.



Same session – but we used the common Pb infested Penglai as the primary and treated 91500 as the unknown – comes out at 1065 Ma.